

# Trends in Plant Science

### Forum

# Incorporation of NPP into forest CH₄ efflux models

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Forest soils are the largest atmospheric methane (CH<sub>4</sub>) sinks in terrestrial ecosystems, but models simulating this uptake have considerable uncertainties. Soil organic matter derived from aboveground vegetation net primary productivity (NPP) significantly influences CH<sub>4</sub> uptake; therefore, we propose that the incorporation of NPP into global CH<sub>4</sub> uptake models will greatly improve model predictions.

### Significance and uncertainty of forest methane sinks

CH<sub>4</sub> is the second most influential greenhouse gas, accounting for up to 25% of the global greenhouse effect. The singlemolecule warming potential of CH<sub>4</sub> is 25-30 times that of carbon dioxide (CO<sub>2</sub>) [1]. Consequently, small changes in the concentration of CH<sub>4</sub> in the atmosphere can have a disproportionally large impact on global warming [2]. Given that atmospheric CH<sub>4</sub> concentrations have increased by a factor of ~2.5 since the Industrial Revolution (722 parts per billion to 1866 parts per billion in 2019), there is considerable interest in tracking atmospheric CH<sub>4</sub> fluxes (e.g., FLUXNET-CH<sub>4</sub>) [1]. There are many CH<sub>4</sub> sources (e.g., wetlands, biomass burning, landfills, etc. [3]), but only two sinks: (i) hydroxyl radicals in the troposphere reacting with CH<sub>4</sub> to produce carbon monoxide and CO<sub>2</sub>, thereby absorbing ~90% of atmospheric  $CH_4$ ; and (ii) upland terrestrial soils, which provide a niche for microorganisms that oxidize  $\sim 9-10\%$  of atmospheric CH<sub>4</sub> [2].

Forests are the largest atmospheric  $CH_4$  sinks among the terrestrial ecosystems, comprising 52% of the total terrestrial  $CH_4$  sink [2,4].

Models predicting the extent of CH<sub>4</sub> oxidation in forested areas currently produce estimates that can differ substantially when compared with both each other and the observed data [5]. This inconsistency both between models and with observed data is potentially driven by the impact of localized or transitory extreme climate events (e.g., forest fires and drought) that cannot be reliably simulated [2,6], and by the spatial variability of various environmental factors that influence the relative oxidative capacity of a forest soil, confounding extrapolation of local studies to wider scales [4,6]. More importantly, our understanding of the microbial processes that underpin forest CH<sub>4</sub> sink capacity remains incomplete, and the impact of potentially important factors, such as the availability of organic matter to support methanogenesis, is not accounted for [7,8]. Consequently, models simulating the rates and extent of forest CH<sub>4</sub> sinks have limited accuracy, and this uncertainty has substantial flow-on effects for efforts to understand global atmospheric CH<sub>4</sub> balances [5,6].

## Microbial processes driving the soil methane balance

The extent of the potential  $CH_4$  sink of a given soil is determined by the balance between two processes:  $CH_4$  production by soil methanogens, which use materials produced during plant decomposition under anaerobic conditions [8], and  $CH_4$ oxidation under aerobic conditions by methanotrophs [7,8] (Figure 1). Substrates for methanogenesis are derived from the decomposition of plant-derived organic matter. Under anaerobic conditions, complex organic matter is converted to monosaccharides and then fermented into fatty acids,  $CO_2$ , and hydrogen gas (H<sub>2</sub>); the fatty acids are then oxidized by syntrophs to produce acetic acid,  $CO_2$ , and  $H_2$ . The  $CO_2$  and  $H_2$  produced throughout this process are used by hydrogenotrophic methanogens to generate CH<sub>4</sub>, whereas acetic acid is used as a substrate by acetoclastic methanogens [9]. Although anaerobic conditions are essential for this pathway to occur, the availability of suitable substrates is also a critical determinant. The oxidation of CH44 is mediated by methanotrophs, which are Gram-negative bacteria within the Verrucomicrobia, Proteobacteria, and novel NC10 phylum [7]. The key step in soil CH<sub>4</sub> oxidation is catalyzed by CH<sub>4</sub> monooxygenase, which converts CH<sub>4</sub> into methanol, which is then further broken down by the bacteria to produce energy [7,8]. Aerobic conditions facilitate gas exchange between the aboveground and soil atmospheres, allowing methanotrophs access to CH<sub>4</sub>.

The balance between CH<sub>4</sub> production and oxidation is dictated by soil moisture because of its impact on gas diffusion through the soil; however, it is also affected by soil temperature and texture [7,8]. In general, forest soils are dry and well aerated, creating an aerobic environment that favors methanotrophs; nevertheless, in situ observations have demonstrated that transient increases in soil moisture content caused by heavy rain can create anaerobic conditions that cause the soil to become a source of CH<sub>4</sub> [6,10]. Other events, such as compaction, can also reduce aeration, creating anaerobic sites. Direct interactions may also be possible, because soil methanotrophs can use the CH<sub>4</sub> produced by methanogens before it leaves the soil atmosphere (Figure 1).

# Factors predicting methane sink capacity in forest soils

To estimate the balance between aerobic and anaerobic conditions, models such as the Methanotrophy Model (MeMo v1.0), focus on climate, soil physical and chemical

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Figure 1. Schematic showing the pathway by which plant productivity can affect soil methane ( $CH_4$ ) efflux via organic matter production in forest ecosystems. The  $CH_4$  efflux across the soil–atmosphere interface is determined by production and oxidation processes in the anaerobic and aerobic soil sites, with soil organic matter providing the resources required for methanogenic activity.

properties, the diffusion of  $CH_4$  in the soil, and the kinetics of  $CH_4$  monooxygenase as the key drivers of  $CH_4$  dynamics [5]. However, the impact of available organic matter on  $CH_4$  production is not as thoroughly considered. Given the fundamental importance of soil organic matter as the raw material for methanogenic activity [9], we suggest that the NPP of the aboveground vegetation should also be integrated in models, such as MeMo, to improve estimates of forest  $CH_4$  sink size [4–6].

This viewpoint is supported by consideration of the fate of carbon captured through plant growth. It has been established that 30-60% of photosynthetically fixed carbon is allocated to plant roots, more than half of which is eventually transferred to the soil in various forms [11]; the aboveground plant litter adds further to this pool. Under anaerobic conditions, a large proportion of this carbon is rapidly transformed and emitted as  $CH_4$  [9,12]. In a recent wetland  $CH_4$ emission model, the researchers explored NPP as a predictor and found that it explained variations in CH<sub>4</sub> flux better than did established climatic factors, improving the overall accuracy of model predictions [13]. On the basis of this outcome, it is reasonable to consider that, with greater

NPP from a given soil, the pool of organic matter available to support methanogenesis is also proportionally greater. Therefore, when conditions are suitable for methanogens to become active, the rate of methanogenesis will be greater than for soils in which organic matter is less available. If the availability of organic matter becomes a rate-limiting step in methanogenesis, this could lead to variations in the CH<sub>4</sub> balance that can be predicted by NPP.

## Concluding remarks and perspective

Direct measurement of NPP at the level of precision required to support improvements to broad-scale models of the forest soil CH<sub>4</sub> balance are unlikely, but the advent of remote-sensing capabilities provides a new opportunity to develop and apply useful proxies. For example, calculations of vegetation index values from remotely captured spectral characteristics have been used to support the normalized difference vegetation index (NDVI), which has become an established tool for modeling NPP effectively [14]. This offers the potential to generate NPP estimates rapidly and accurately at scalable levels of spatial resolution for forest ecosystems and, thus, explore the relationships among forest NPP, NDVI, and CH<sub>4</sub> balance at the ecosystem level (Figure 2). Examination of these relationships will confirm the potential value of integrating NPP values into forest CH<sub>4</sub> sink models and verify whether they can provide additional parameters to improve the accuracy and, therefore, the utility of these models.

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Figure 2. Schematic showing the relationships between forest net primary productivity (NPP), normalized difference vegetation index (NDVI), and methane ( $CH_4$ ) production potential at the ecosystem scale. We can use NDVI based on remote sensing to estimate the NPP of forest ecosystems: when forest NPP is high, the soil  $CH_4$  production potential is higher and vice versa. Abbreviation: SOM, soil organic matter.

### **Declaration of interests**

No interests are declared.

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